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Turbulence and Convection Parameterizations: the Eddy-Diffusivity/Mass-Flux (EDMF) Approach and its Implementation into the GFS Model

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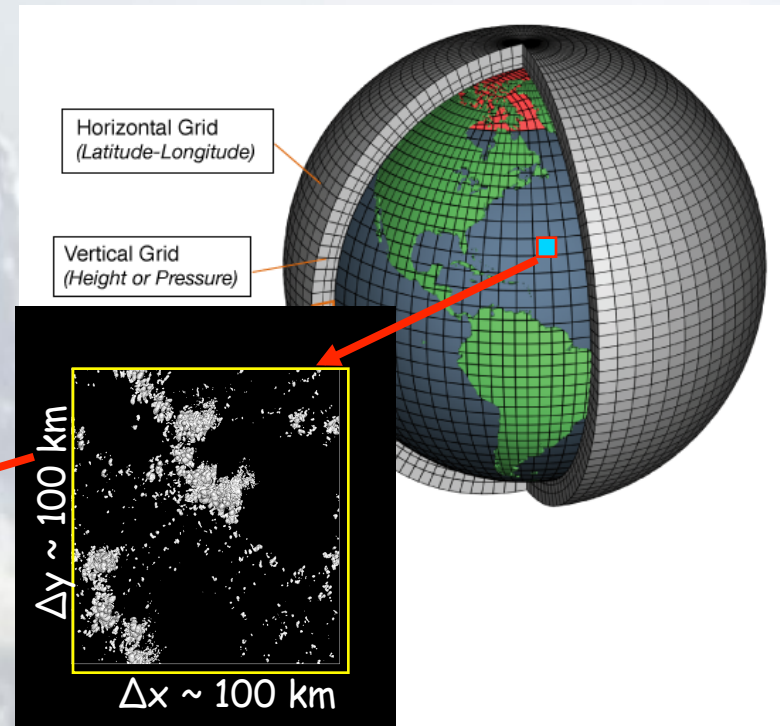
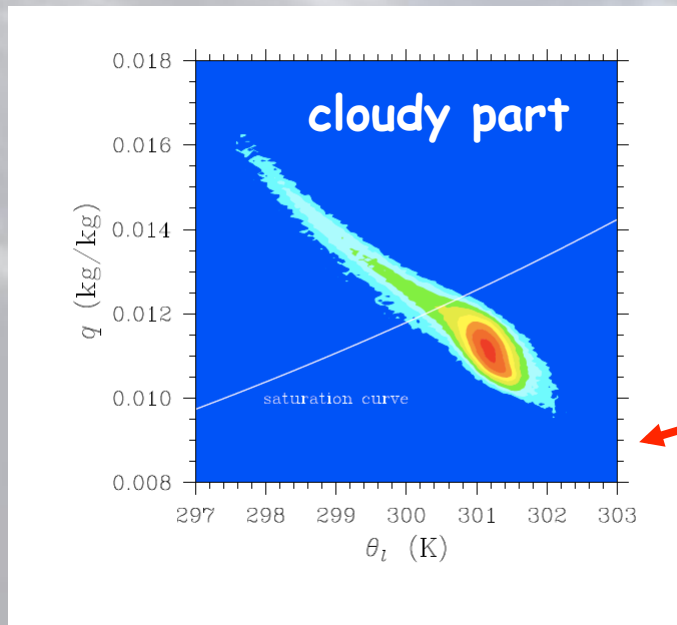
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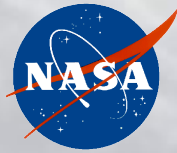
Turbulence and Convection in Weather and Climate Models



$$\varphi = \overline{\varphi} + \varphi' \quad \rightarrow$$

$$\frac{\partial \overline{\varphi}}{\partial t} + \frac{\partial}{\partial x} (\overline{u\varphi}) + \frac{\partial}{\partial y} (\overline{v\varphi}) + \frac{\partial}{\partial z} (\overline{w\varphi}) = - \frac{\partial}{\partial z} (\overline{w'\varphi'}) + \overline{S},$$

Key issue: Turbulence and convection parameterizations



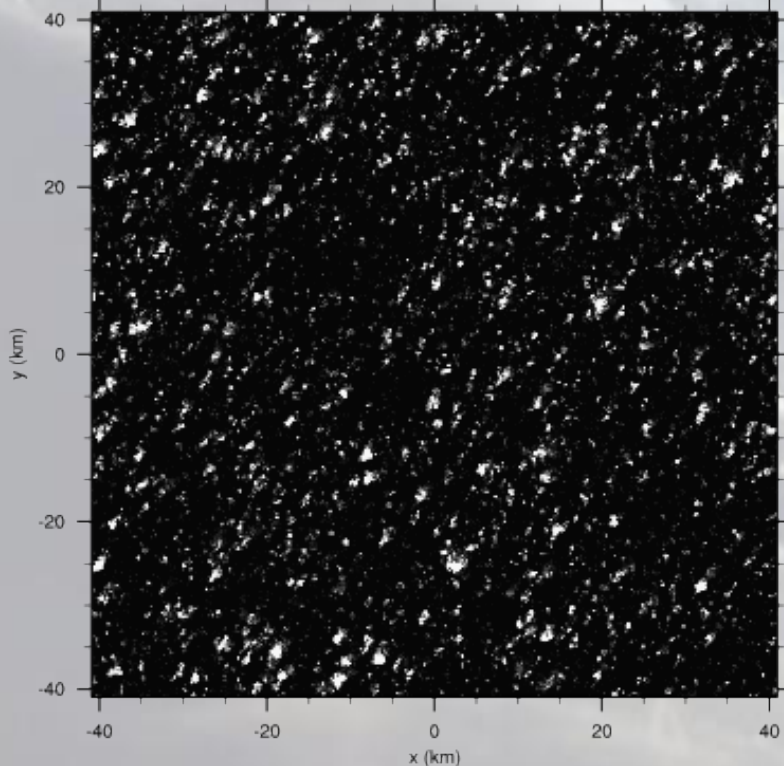
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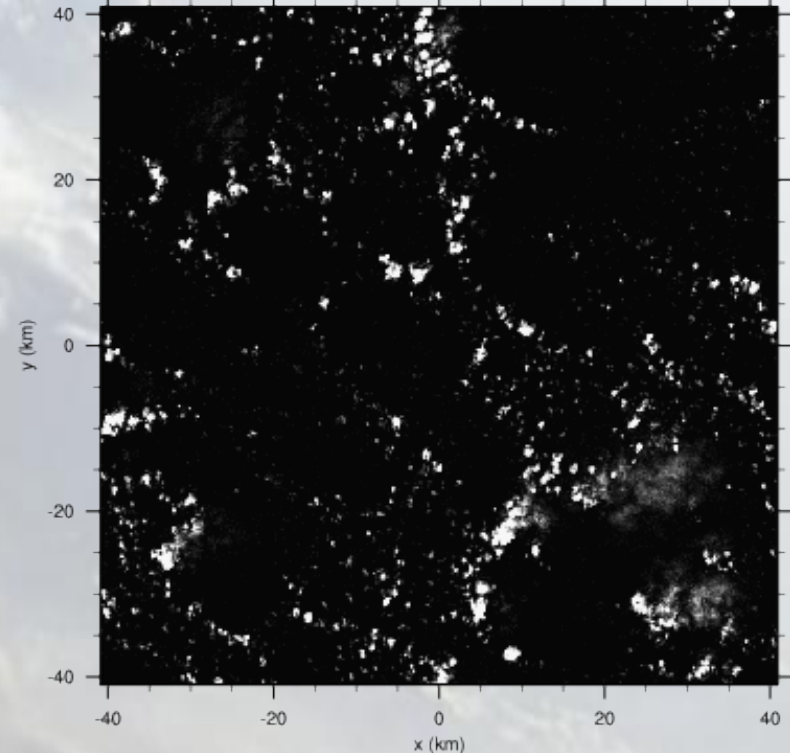
Large-Eddy Simulation (LES) models

- LES models solve filtered version of Navier-Stokes equations
- High-resolutions ($\sim 10\text{-}100\text{m}$) in all 3 dimensions
- LES models resolve most of the essential turbulence/convection
- Closures still needed for scales $< 10\text{m}$ (but simpler to do)

Clouds not allowed to rain



Clouds allowed to rain



Matheou et al., MWR, 2011



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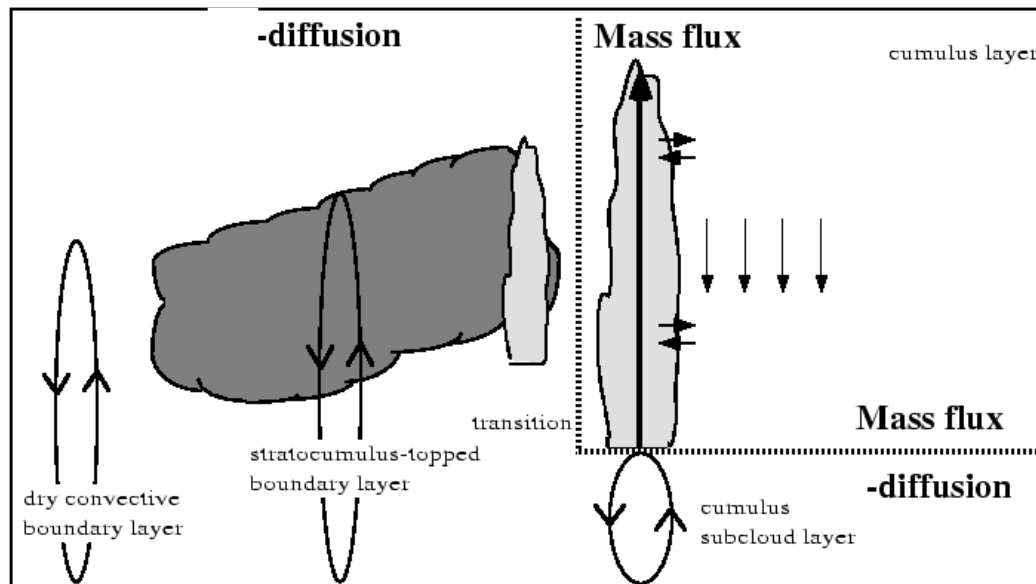
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Clear and Cloudy Convective Boundary layers: Need for Unified/Integrated Approaches

$$\overline{w' \phi'} \cong -K \frac{\partial \bar{\phi}}{\partial z}$$

$$\overline{w' \phi'} \cong M(\phi_u - \bar{\phi})$$

$$\frac{\partial \bar{\phi}}{\partial t} \cong -\frac{\partial}{\partial z} (\overline{w' \phi'}) + \bar{S}$$



Courtesy de Roode & Siebesma

Modularity leads to problems:

- “Double counting” of processes
- Interface problems
- Problems with transitions between different regimes

Key Problem: artificial modularity in vertical mixing parameterizations



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Unified Approach: Eddy-Diffusivity/Mass-Flux (EDMF)

Dividing a grid square in two regions (updraft and environment) and using Reynolds decomposition and averaging leads to

$$\overline{w'\varphi'} = a_u \overline{w'\varphi'_u} + (1 - a_u) \overline{w'\varphi'_e} + a_u(1 - a_u)(w_u - w_e)(\varphi_u - \varphi_e)$$

where a_u is the updraft area. Assuming $a_u \ll 1$ and $w_e \sim 0$ leads to

$$\overline{w'\varphi'} = \overline{w'\varphi'_e} + a_u w_u (\varphi_u - \bar{\varphi})$$

ED closure: assuming ED for 1st term and neglecting 2nd term

MF closure: neglecting 1st term and assuming $M = a_u w_u$

EDMF:
$$\overline{w'\varphi'} = -k \frac{\partial \bar{\varphi}}{\partial z} + M(\varphi_u - \bar{\varphi})$$

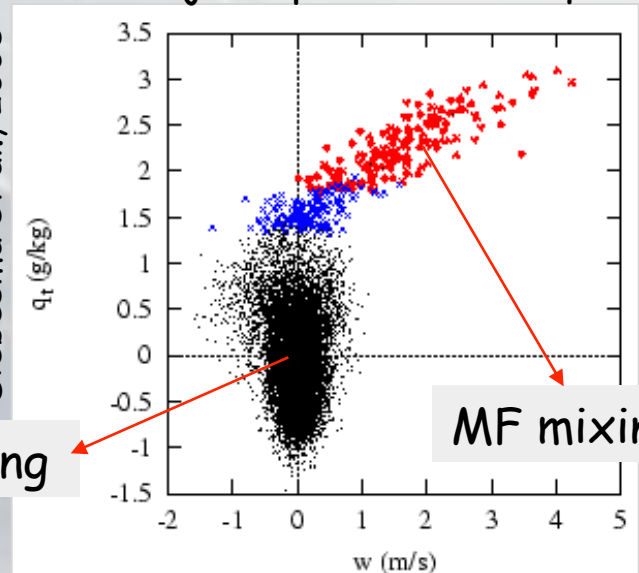
Siebesma & Teixeira, 2000

Bimodal joint pdf of w and q_t

Siebesma et al., 2008

ED mixing

MF mixing



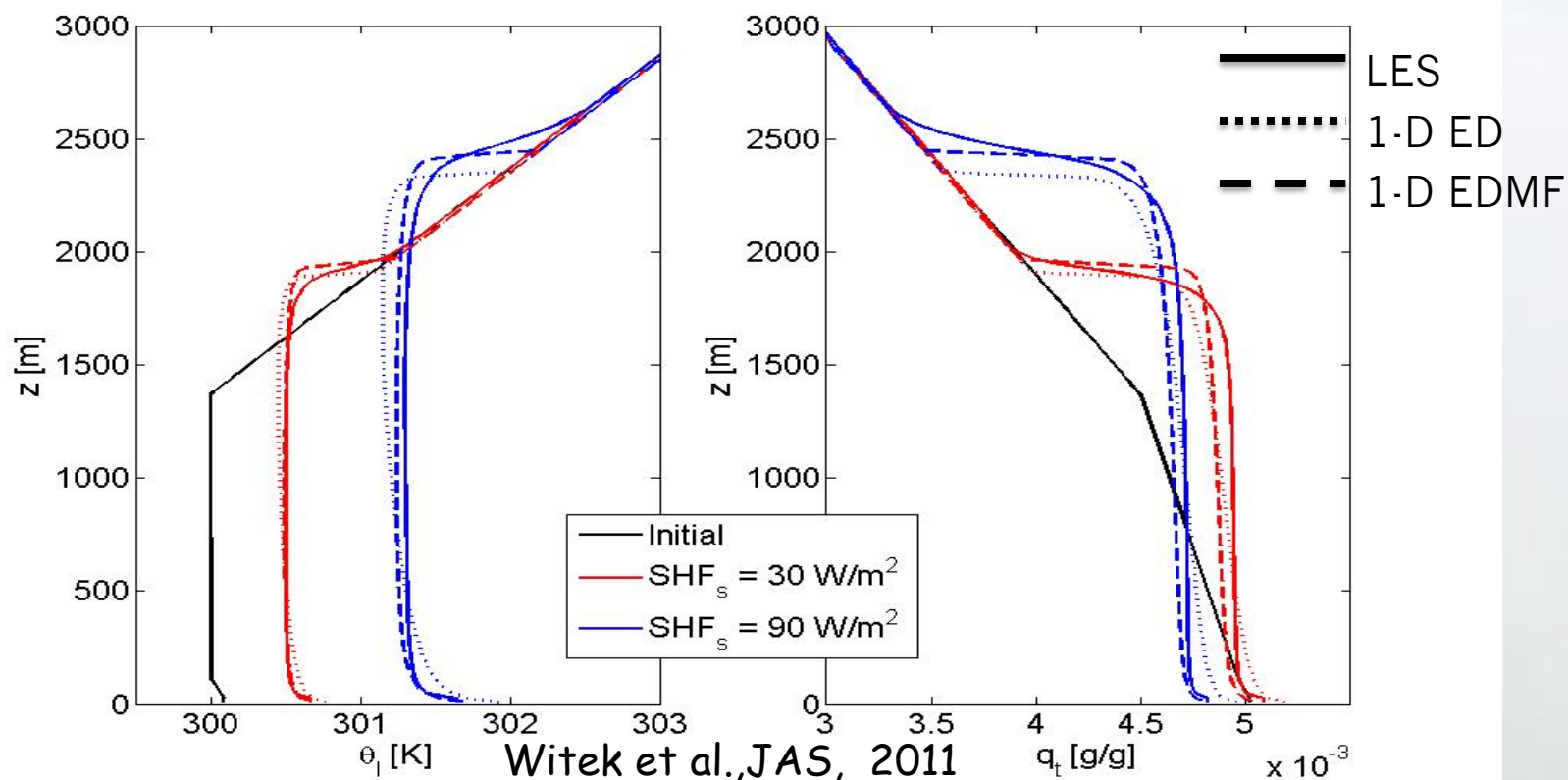
EDMF represents turbulence/convection in an integrated manner



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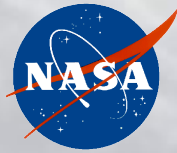
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Dry Convective Boundary Layer: θ and q_t vertical profiles after 6 hours



Dry EDMF simulations:

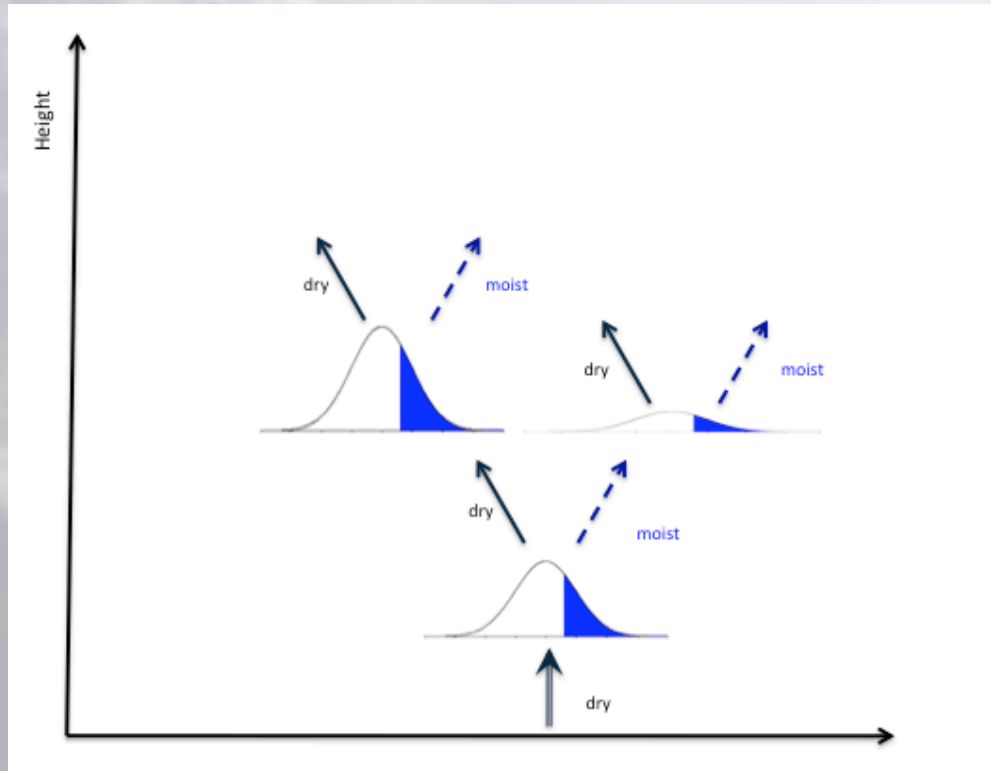
- surface layer more realistic
- neutral profile in the well-mixed layer
- larger entrainment leads to better inversion height
- inversion layer too sharp compared to LES



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Stochastic Plume for moist EDMF: using PDF of updraft properties



Suselj et al., JAS, 2012, 2013

- 1) Estimate PDF of plume/updraft properties (T , q , w)
- 2) Sample PDF to generate a variety of plumes (diff. properties)
- 3) Integrate different plumes in the vertical

Produces more realistic results than purely deterministic parameterization

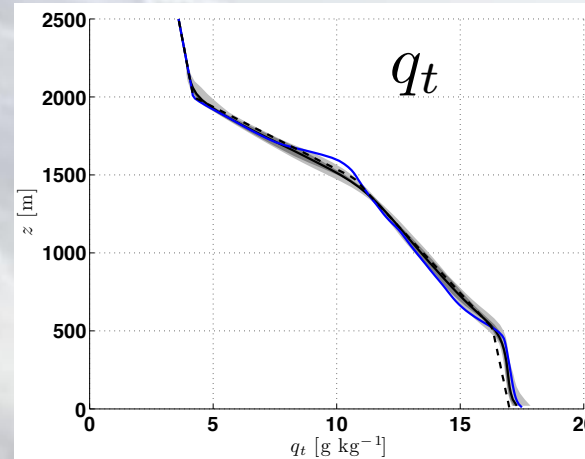
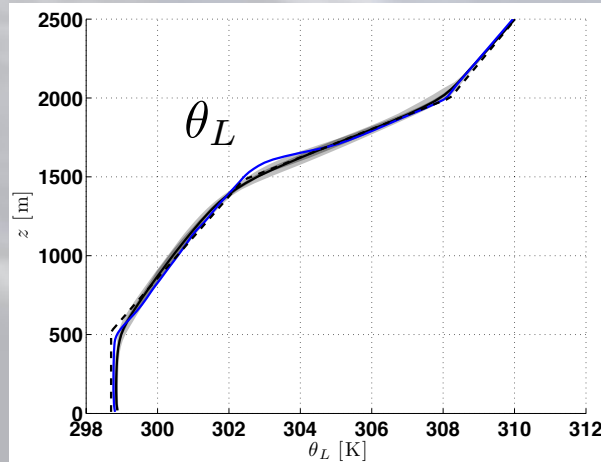


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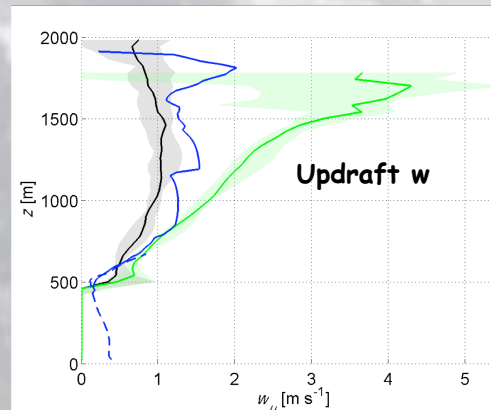
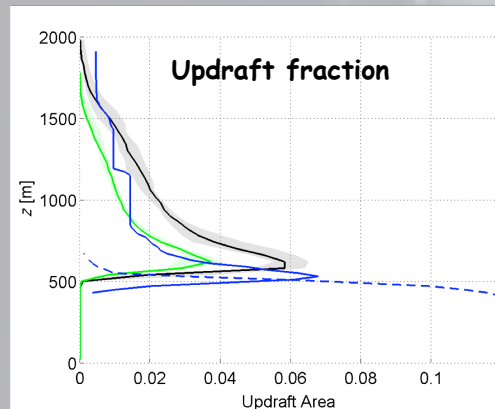
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EDMF simulation of shallow cumulus BOMEX case: comparison with LES

Mean profiles between 3rd and 4th simulation hour



Single column model
LES, mean



Single column model, dry
Single column model, moist
LES, cloud core, mean
LES, cloud core, range
LES, clouds, mean
LES, clouds, range

Suselj et al, JAS, 2012, 2013

New aspect: Using PDF of updraft properties



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Dry EDMF Implementation into GFS

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Vertical diffusion equation

Reference

$$\frac{\partial C}{\partial t} = \frac{\partial}{\partial z} \left[K_c \left(\frac{\partial C}{\partial z} - \gamma_c \right) \right]$$

$$C \in (\theta, q_t, u, v)$$

K - diffusion coefficient

$$\gamma_c = b \frac{\overline{w'c'^s}}{w_s} - \text{countergradient term}$$

$$w_s \equiv \left(u_*^3 + 0.7kw_*^3 \right)^{1/3} - \text{mixed-layer velocity scale}$$

EDMF

$$\frac{\partial C}{\partial t} = \frac{\partial}{\partial z} \left[K_c \frac{\partial C}{\partial z} - M(C_u - C) \right]$$

$M = aw_u$ - mass-flux term

C_u - updraft characteristics

$w_u(z), \theta_u(z), q_{t,u}(z)$ - unknown variables

$a \approx 0.1$ - fixed updraft fraction



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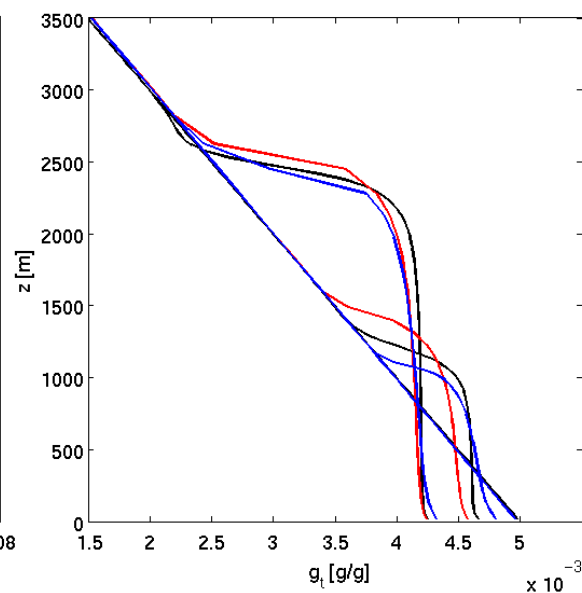
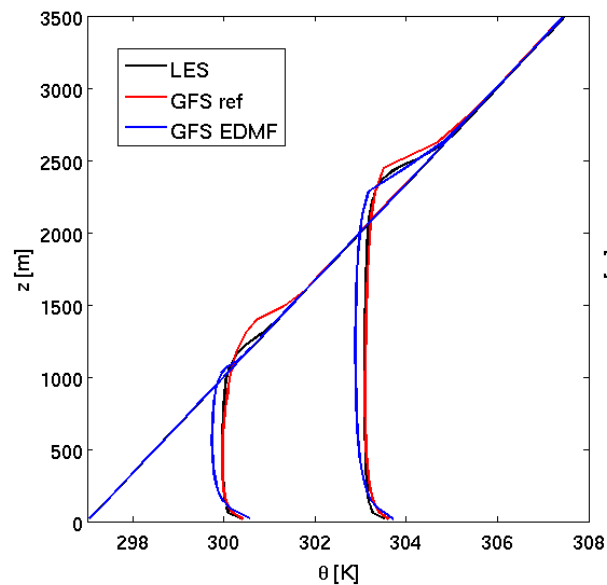
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Dry EDMF in GFS: SCM dry convection simulations

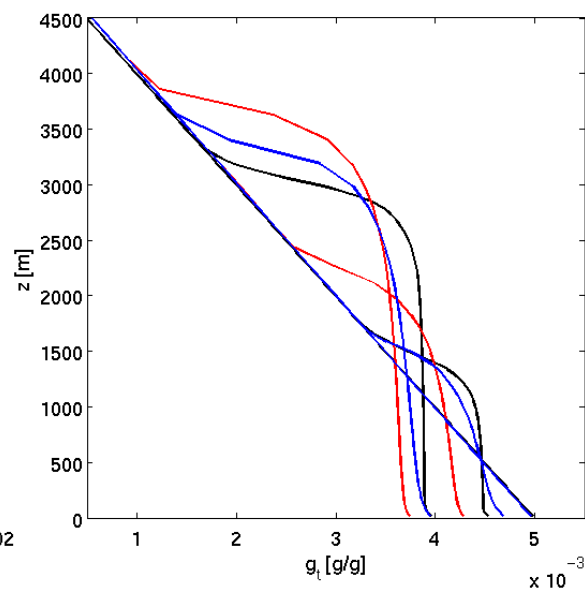
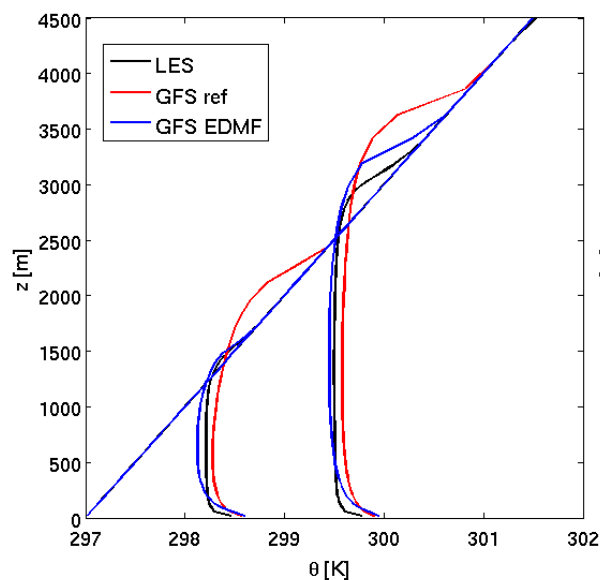


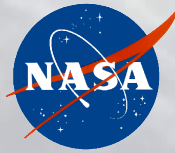
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$\text{SHF}_s = 0.20$
 $\text{grad}(\theta) = 3$



$\text{SHF}_s = 0.10$
 $\text{grad}(\theta) = 1$

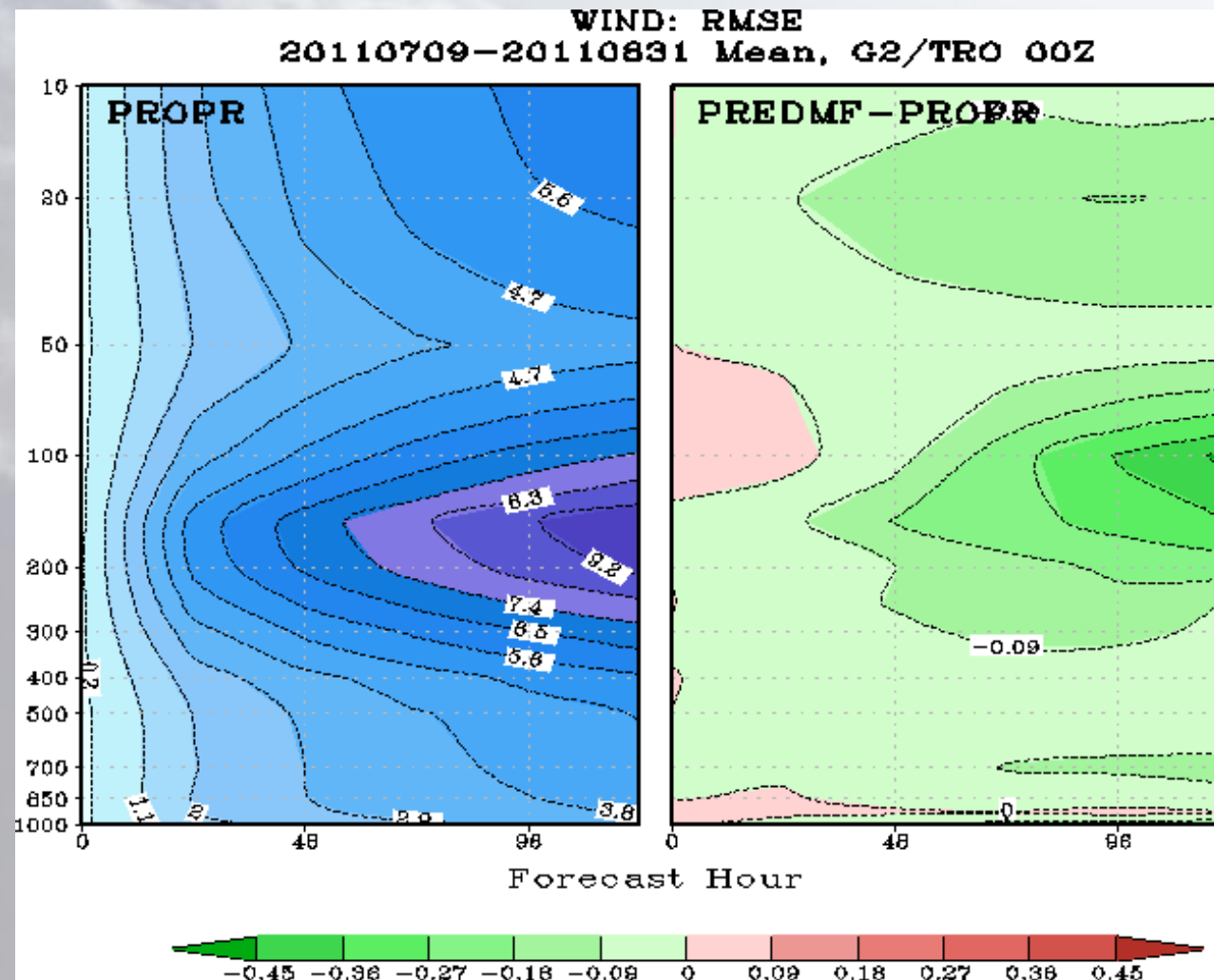




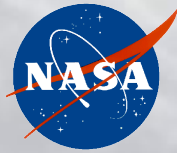
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'Dry' EDMF Implemented into GFS: Data Assimilation experiments



- Neutral Z500 anomaly correlation, but significantly reduced wind errors

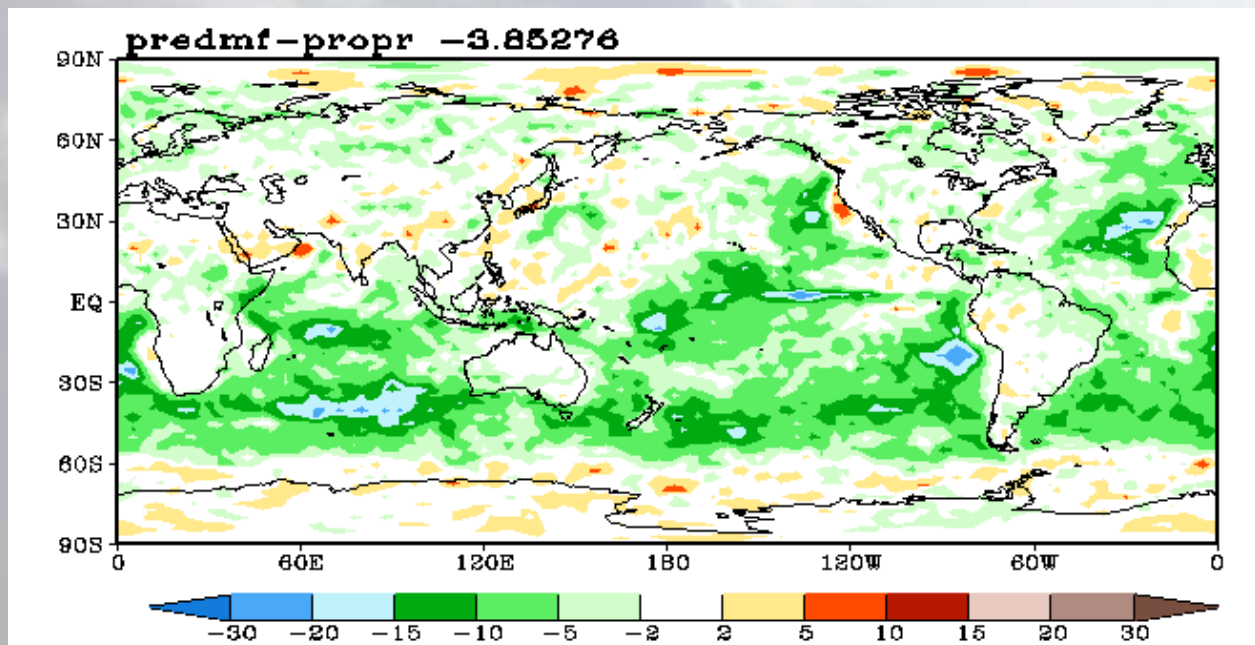


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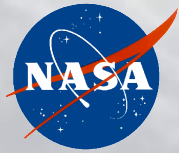
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Dry EDMF in GFS: Clouds

Dry EDMF decreases cloud cover in GFS by 4%



Dry EDMF is just an initial step to a fully moist implementation of EDMF proposed for Phase 2 of our CPT



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Summary

- Realistic parameterization of PBL turbulence and convection is essential for weather and climate prediction
- Modularity problem in models: the need for unified schemes
- New approach: combining Eddy-Diffusivity and Mass-Flux (EDMF)
- Dry convection more realistic with EDMF
- Stochastic EDMF more realistic for moist convection
- Dry EDMF implemented in GFS: neutral Z500, less realistic clouds, substantial reduction of wind errors